

# Mineral indicators of potassium activity in the upper mantle: a review

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# MANTLE METASOMATISM

compositional changes in the mantle wall-rock due to interaction with mantle fluids (Lloyd & Bailey, 1973)

**Cryptic  
metasomatism**

(Dawson, 1984): changes in composition (especially, trace elements) without addition of new phases

**Stealth  
metasomatism**

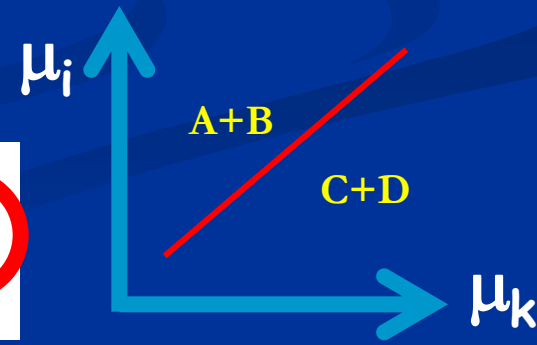
(O'Reilly & Griffin, 2013): addition of new phases, which are indistinguishable mineralogically from the common peridotite phases

**Modal (patent)  
metasomatism**

(Harte, 1983): “the presence of minerals additional to those seen commonly in peridotites”

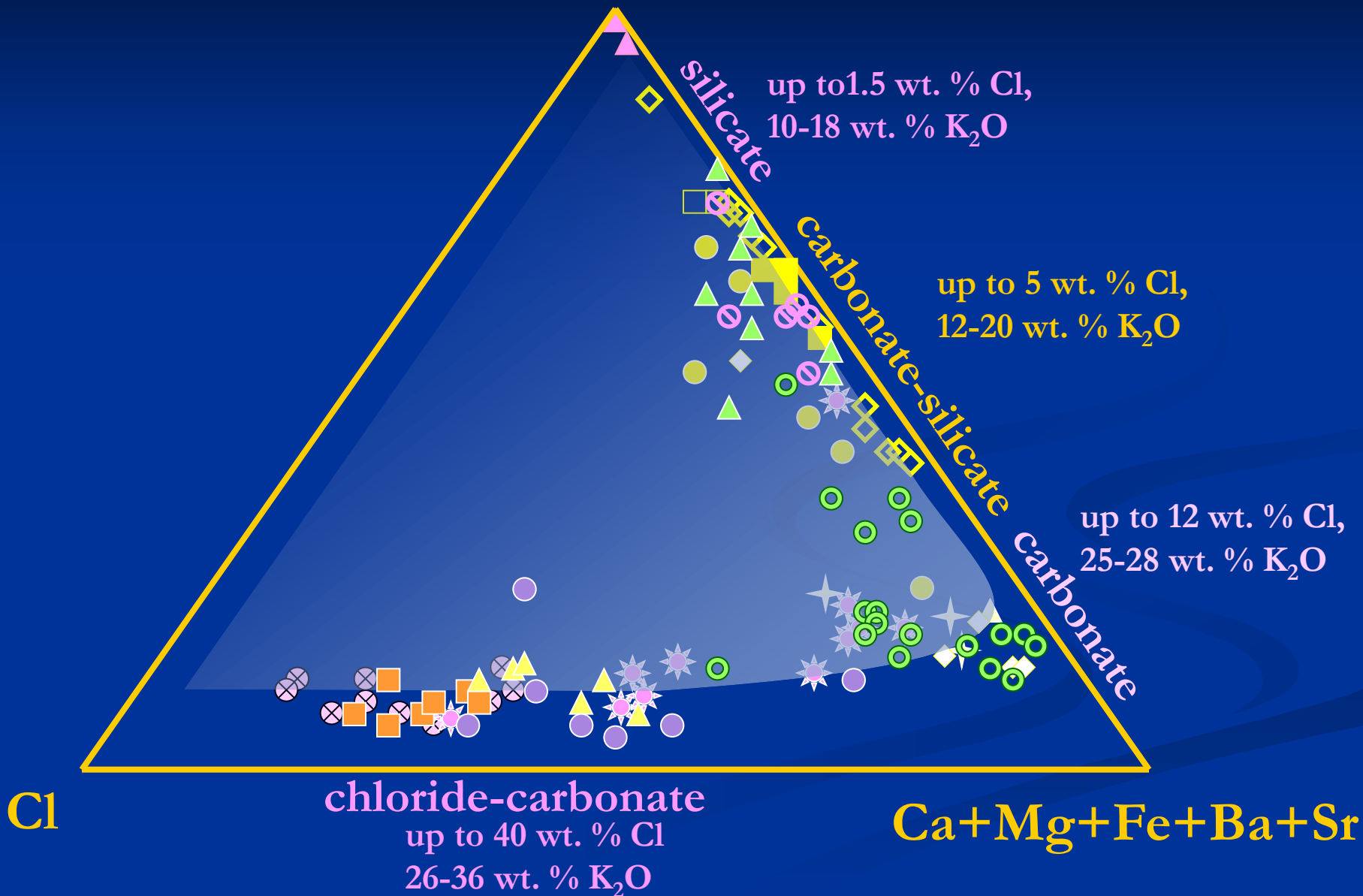
$$dG = -SdT + Vdp + \sum_j^l \mu_j dm_j + \sum_i^k n_i d\mu_i$$

perfectly mobile components

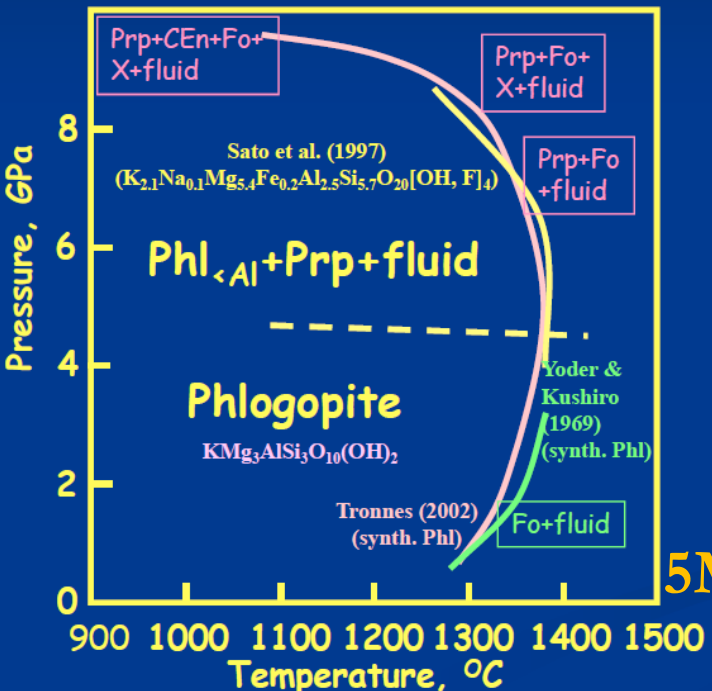
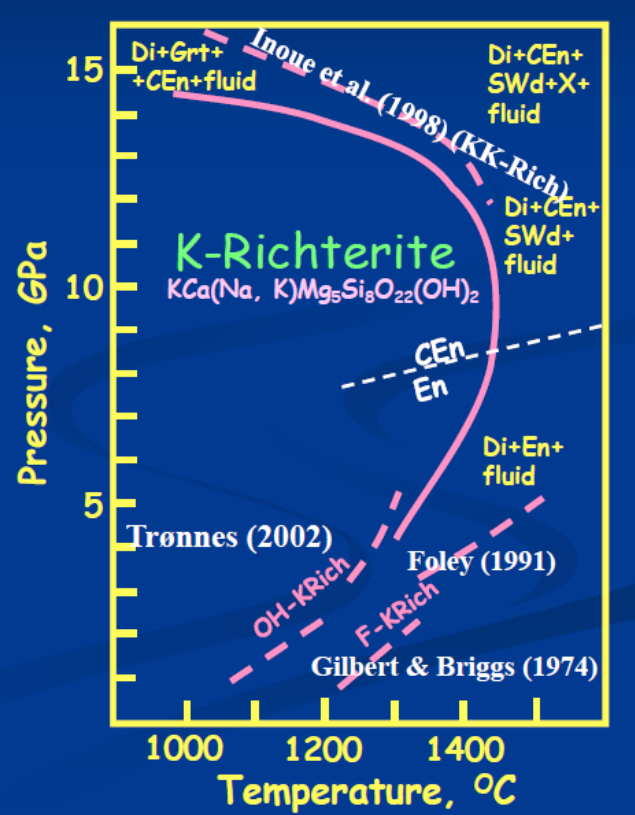
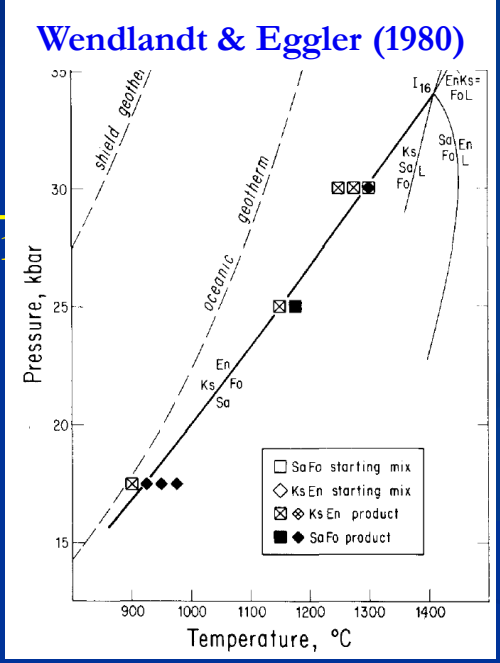
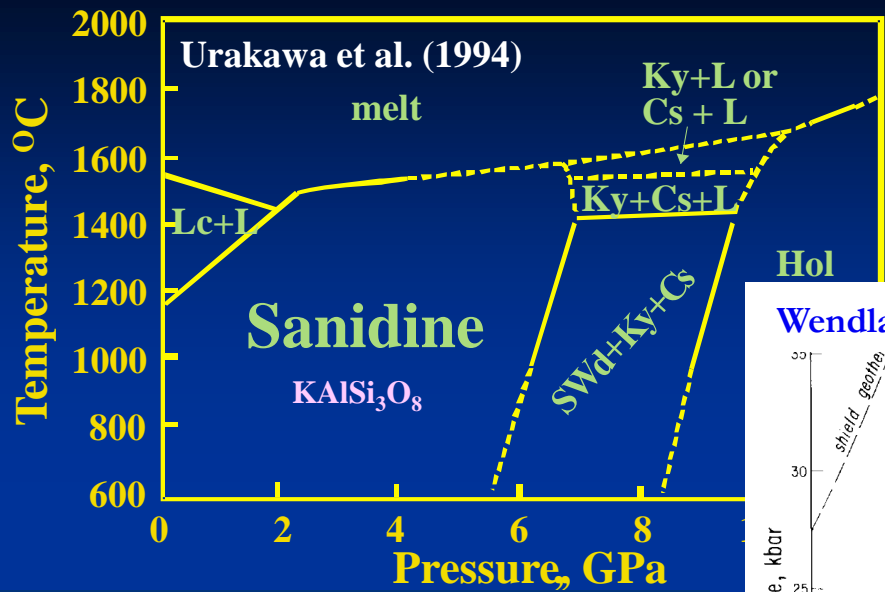


# Alkaline fluids in the upper mantle

Si+Al+Ti+Cr+P



# K-bearing minerals in the upper mantle



**Origin of Phlogopite and Potassic Richterite Bearing Peridotite Xenoliths from South Africa**  
 Ken-ichiro Aoki  
 Contrib. Mineral. Petrol. 53, 145–156 (1975)



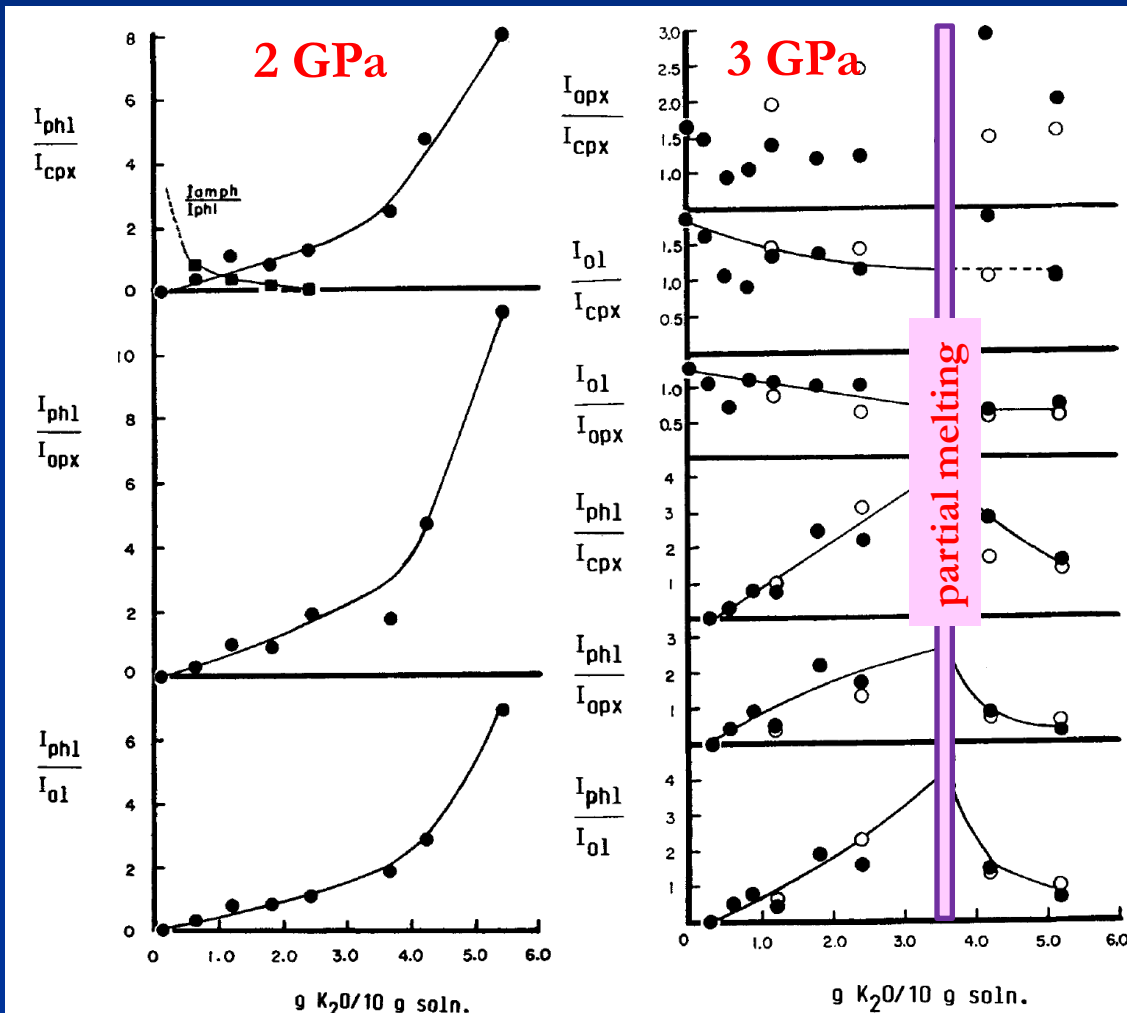
# Pyrolite-K<sub>2</sub>CO<sub>3</sub>-H<sub>2</sub>O system

Patent mantle-metasomatism: Inferences based on experimental studies

Proc. Indian Acad. Sci. (Earth Planet. Sci.), Vol. 99, No. 1, March 1990, pp. 21-37.

Y THIBAUT and A D EDGAR

2 – 3 GPa, 850-950°C



With the increase of the K content of the fluid

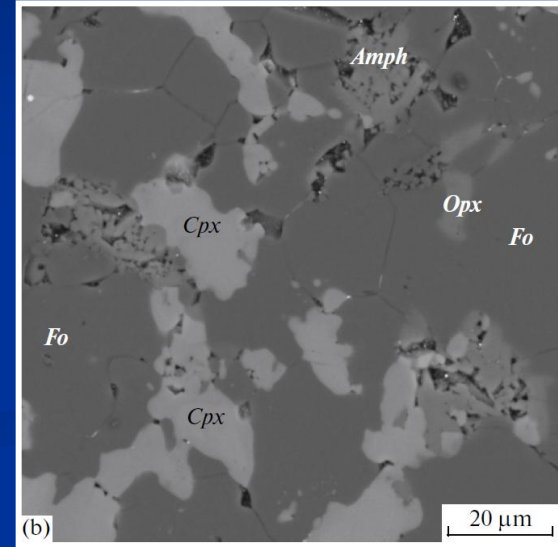
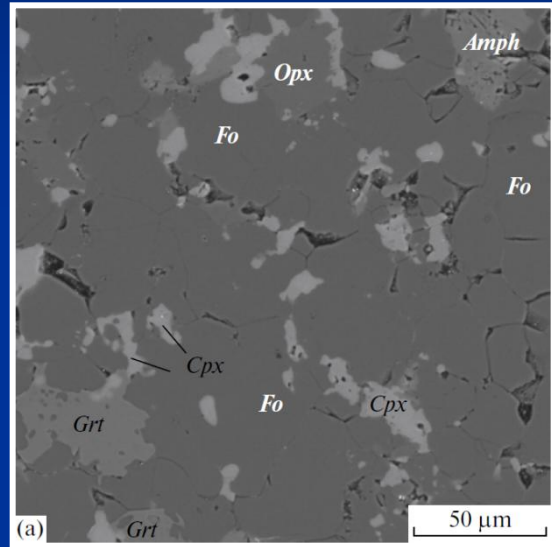
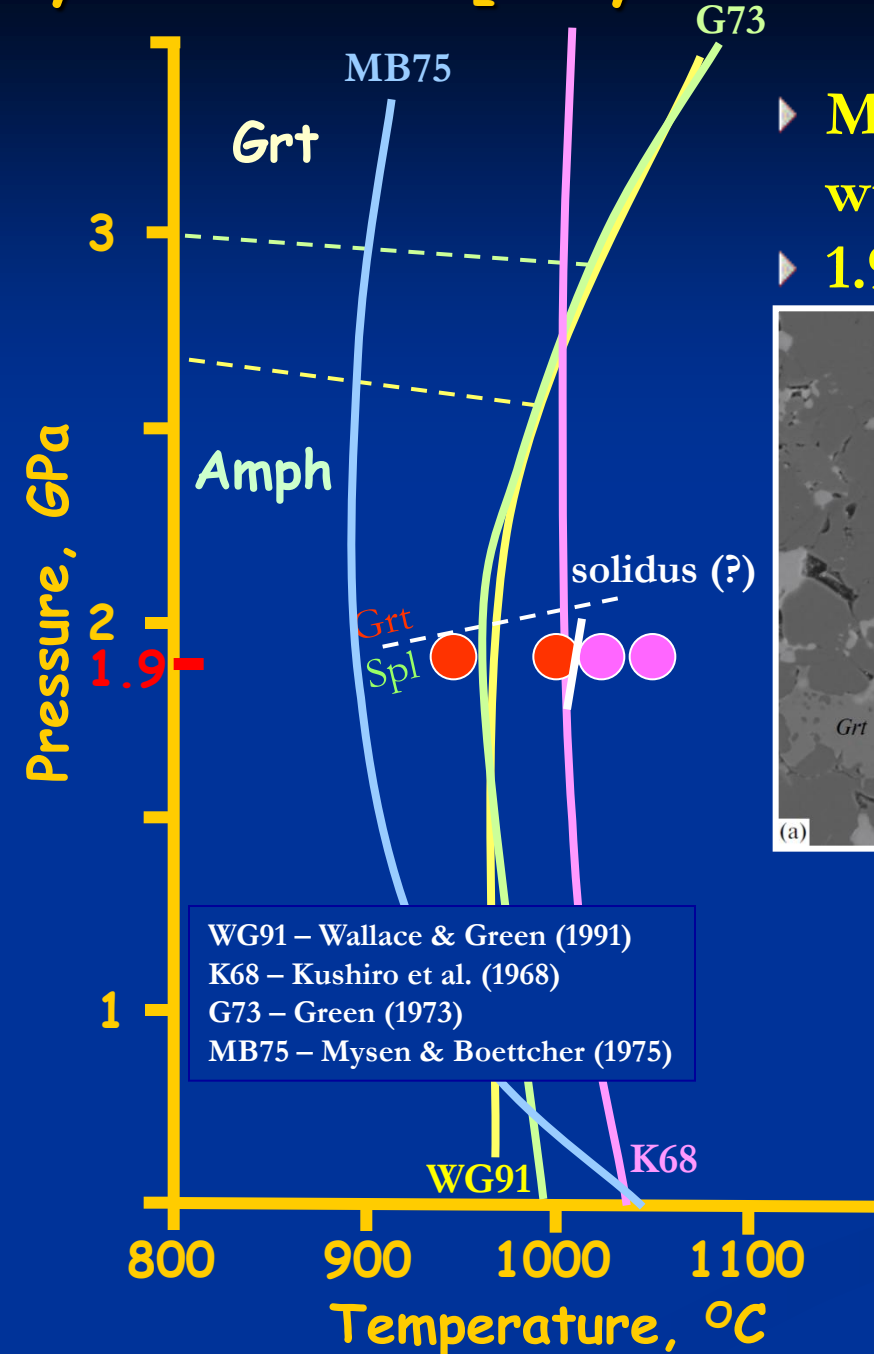
- ✓ Amount of phlogopite systematically increases relative to Ol, Opx and Cpx (decrease at 3 GPa is related to the beginning of partial melting).
- ✓ Al for phlogopite is provided by garnet.
- ✓ Amount of amphibole rapidly decreases at 2 GPa.
- ✓ Increase of the Cpx/Opx and Ol/Opx ratios at 2 GPa.

Patent K-metasomatism at depths 20-30 kbar will rapidly result in the disappearance of garnet, and decline in amounts of spinel, amphibole and orthopyroxene.

# Pyrolite-KCl-H<sub>2</sub>O system

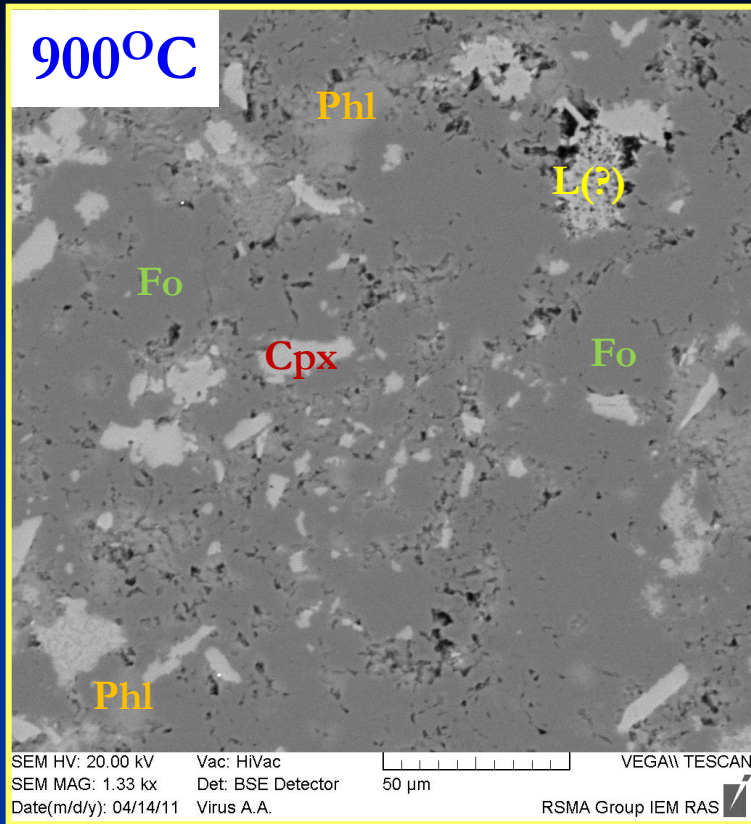
Interaction of Model Peridotite with H<sub>2</sub>O-KCl Fluid:  
 Experiment at 1.9 GPa and Its Implications  
 for Upper Mantle Metasomatism  
*Petrology*, 2013, Vol. 21, No. 6, pp. 599–615.  
 O. G. Safonov<sup>a,b,c</sup> and V. G. Butvina<sup>a</sup>

- ▶ Model peridotite Fo<sub>57</sub>En<sub>17</sub>Prp<sub>14</sub>Di<sub>12</sub> + 0.3 wt. % Na<sub>2</sub>O + 4.4 wt. % H<sub>2</sub>O
- ▶ 1.9 GPa.

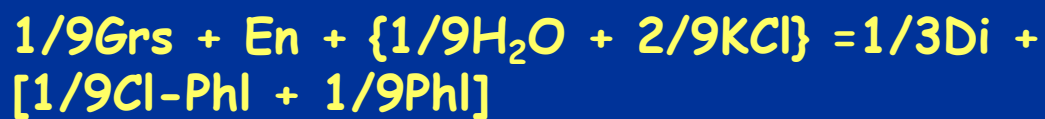
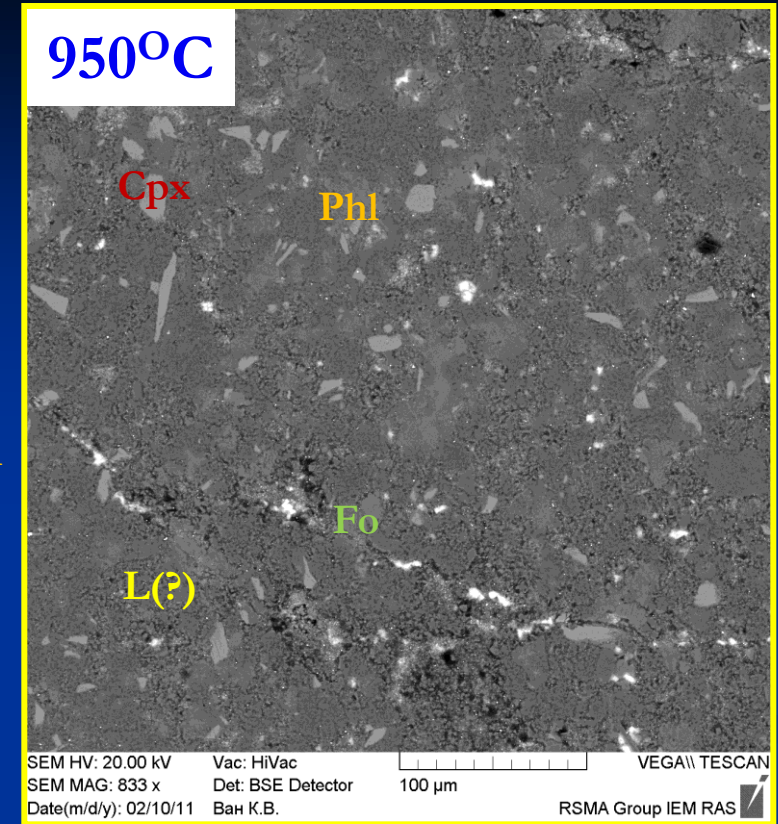


Phase	SiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	MgO	CaO	Na <sub>2</sub> O
Ol	43.48	0.00	56.88	0.05	0.00
Cpx	57.37	0.13	18.94	24.98	0.09
Opx	56.83	6.99	37.74	0.43	0.09
Amph	47.41	15.77	21.04	12.09	2.79
Grt	45.58	25.44	23.81	7.95	0.00

$KCl/(KCl+H_2O) = 0.13$



$KCl/(KCl+H_2O) = 0.39$



- ✓ No Opx, Grt, Amph
- ✓ Leading assemblage Cpx+Fo+Phl
- ✓ Increase of the melt proportion
- ✓ Increase of the Phl content
- ✓ Phl (0.4-0.7 wt. % Cl) is stable above solidus







**Merwinite in diamond from São Luiz, Brazil: A new mineral of the Ca-rich mantle environment**

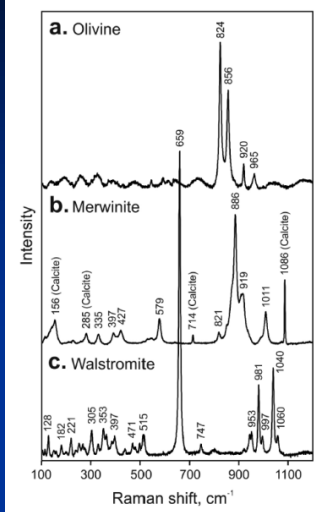
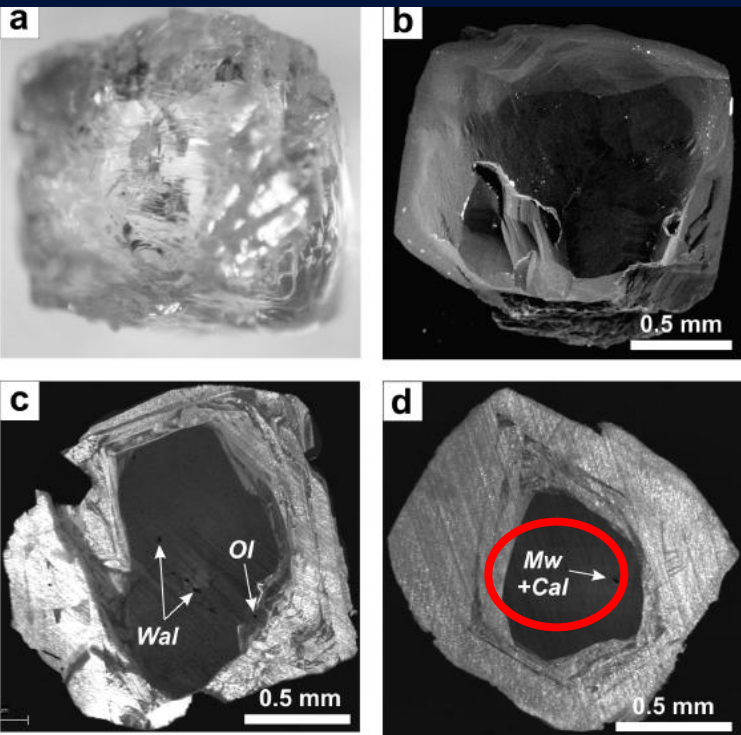
*American Mineralogist, Volume 99, pages 547–550, 2014*

DMITRIY A. ZEDGENIZOV<sup>1,2,\*</sup>, ANTON SHATSKIY<sup>1,2</sup>, ALEXEY L. RAGOZIN<sup>1,2</sup>, HIROYUKI KAGI<sup>3</sup> AND VLADISLAV S. SHATSKIY<sup>1,4</sup>

**Merwinite-structured phases as a potential host of alkalis in the upper mantle**

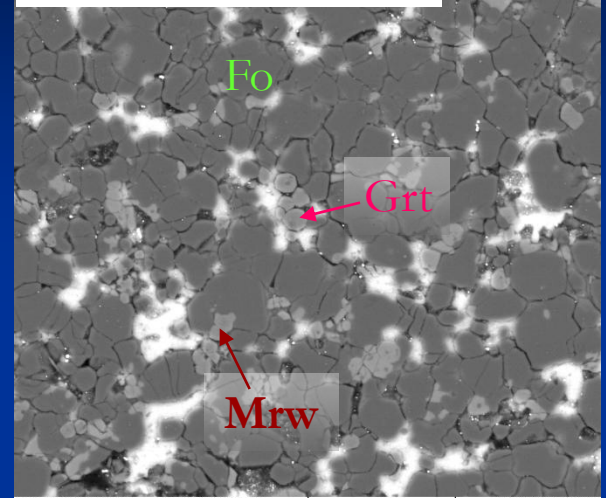
Contrib Mineral Petrol (2015) 170:14

Luca Bindi<sup>1,2</sup> · Oleg G. Safonov<sup>3,4,5</sup> · Dmitriy A. Zedgenizov<sup>6,7</sup>



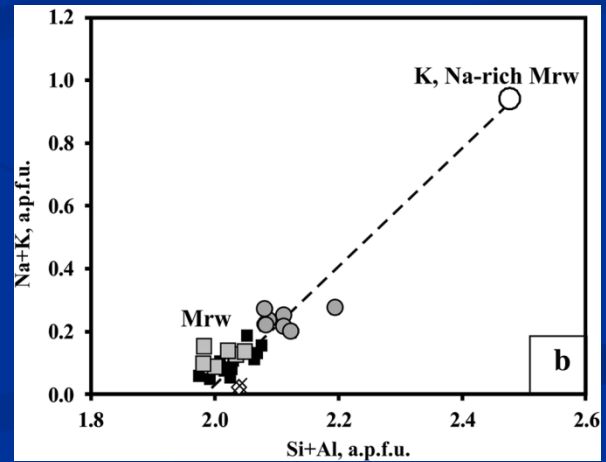
Wal, Ol<sub>86</sub>,  
Cal

7 GPa, 1370°C



SEM HV: 20.00 kV WD: 25.0000 mm VEGA\\ TESCAN  
SEM MAG: 1.33 kx Det: BSE Detector 50 µm  
Date(m/d/y): 11/11/09 Kovalsky RSMA Group IEM RAS

	Walstromite		Merwinite		Olivine	
	wt%	mol%	wt%	mol%	wt%	mol%
SiO <sub>2</sub>	51.6	50.3	36.8	34.2	39.8	33.4
TiO <sub>2</sub>	0.04	0.03	0.02	0.01	0.0	0.0
Al <sub>2</sub> O <sub>3</sub>	0.01	0.01	0.10	0.05	0.0	0.0
Cr <sub>2</sub> O <sub>3</sub>	0.0	0.0	0.02	0.01	0.09	0.03
FeO	0.31	0.26	2.32	1.80	13.6	9.62
MnO	0.00	0.00	0.04	0.03	0.04	0.03
MgO	0.00	0.00	11.7	16.1	45.4	56.8
CaO	47.3	49.4	47.9	47.6	0.12	0.10
Na <sub>2</sub> O	0.03	0.03	0.24	0.22	0.0	0.0
K <sub>2</sub> O	0.00	0.00	0.01	0.00	0.0	0.0
NiO	n.d.	n.d.	n.d.	n.d.	0.25	0.17
Total	99.3	100.0	99.1	100.0	99.3	100.0



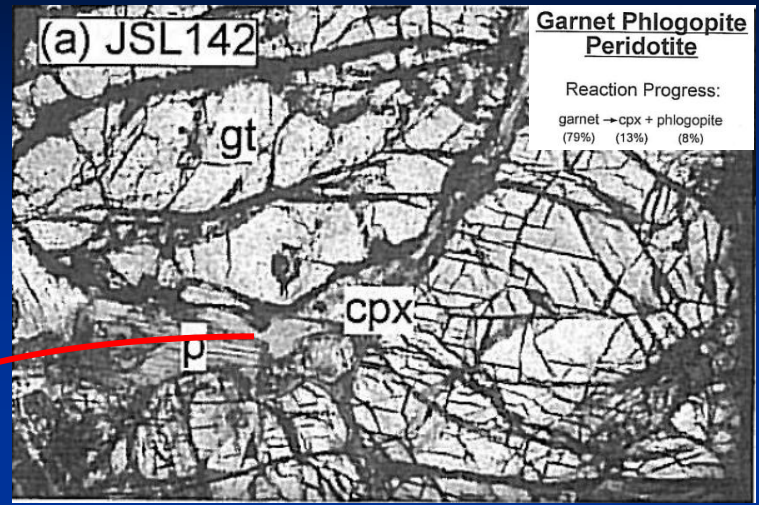
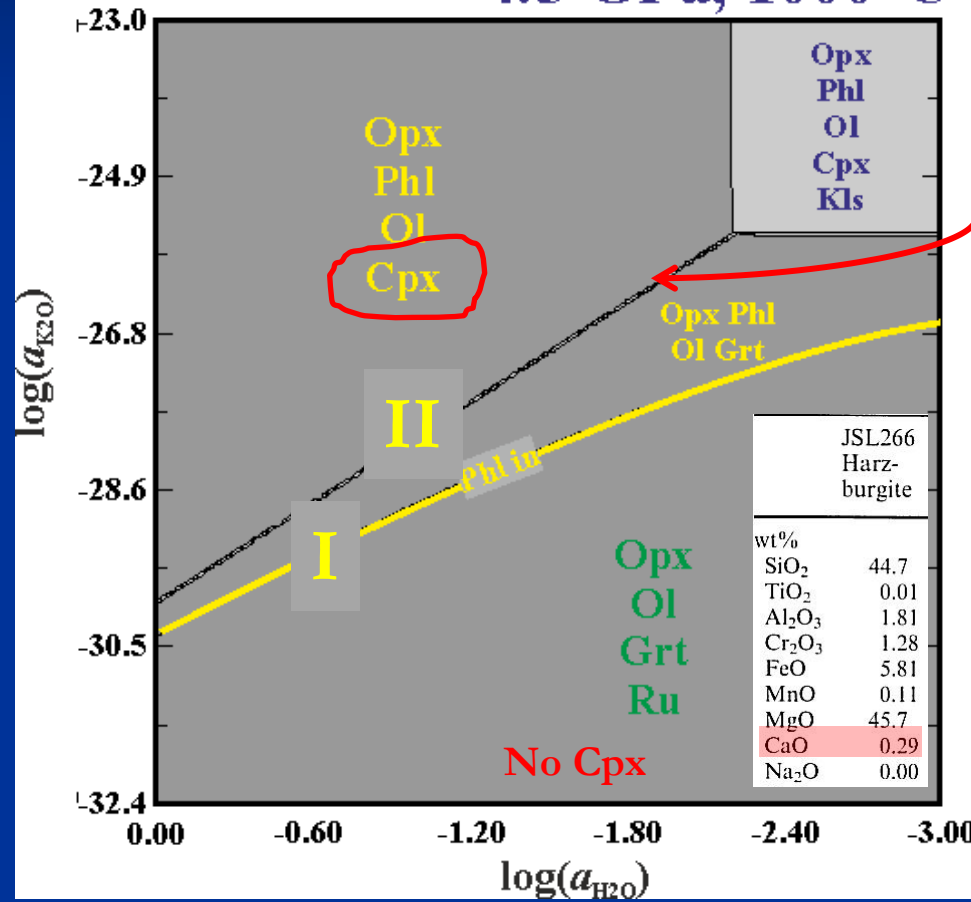
# $\mu_{H_2O} - \mu_{K_2O}$ relations for harzburgite

Esmé van Achterbergh · William L. Griffin  
Johann Stiefenhofer

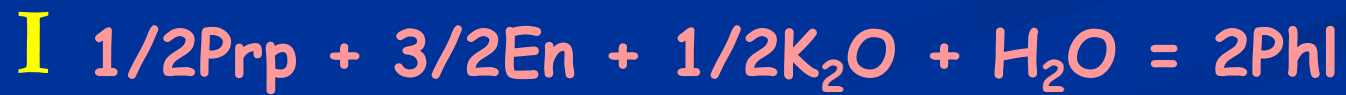
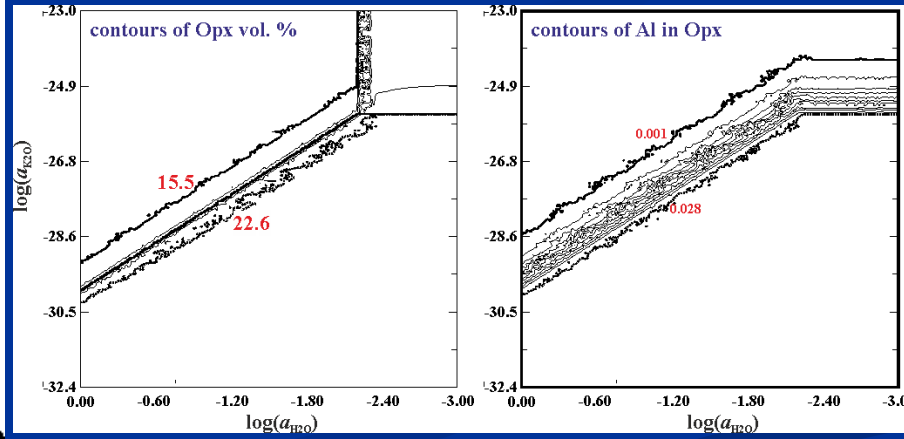
Contrib Mineral Petrol (2001) 141: 397–414

Metasomatism in mantle xenoliths from the Lethakane kimberlites: estimation of element fluxes

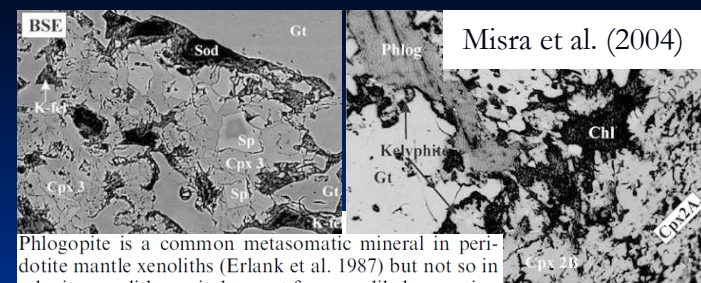
4.5 GPa, 1000°C



**Garnet Phlogopite Peridotite**  
Reaction Progress:  
garnet → cpx + phlogopite  
(79%) (13%) (8%)

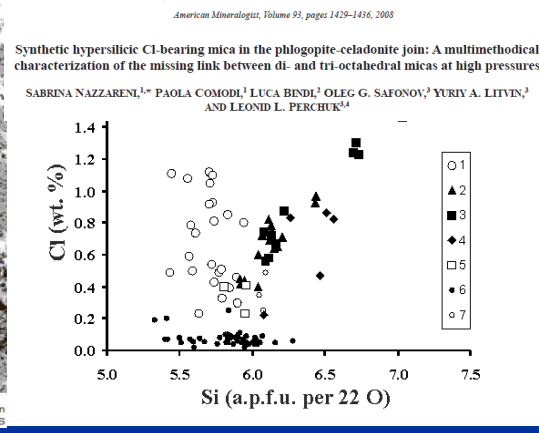
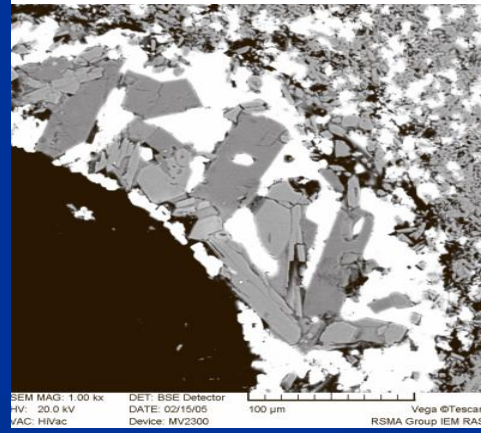
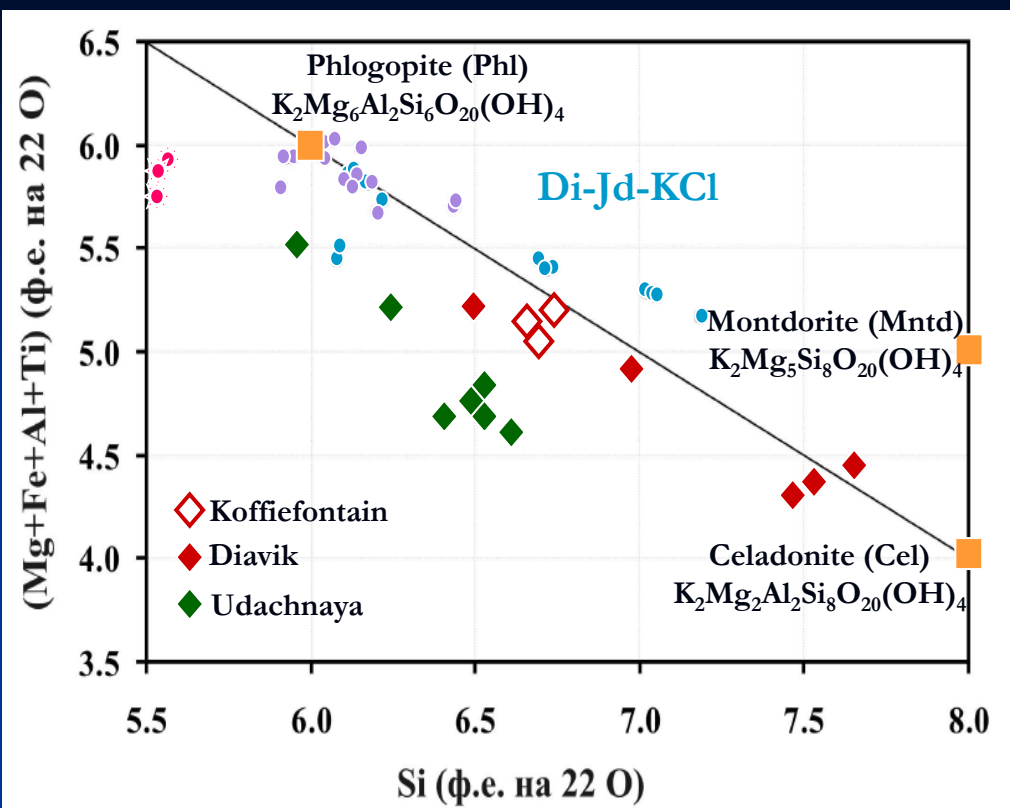
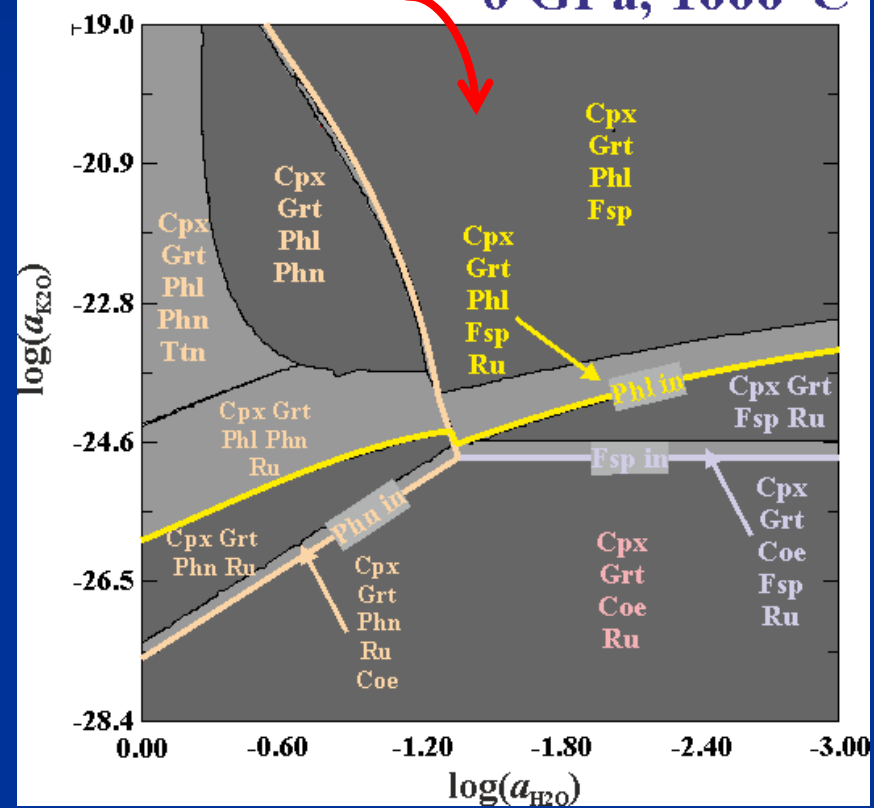


# $\mu_{H_2O} - \mu_{K_2O}$ relations for eclogite

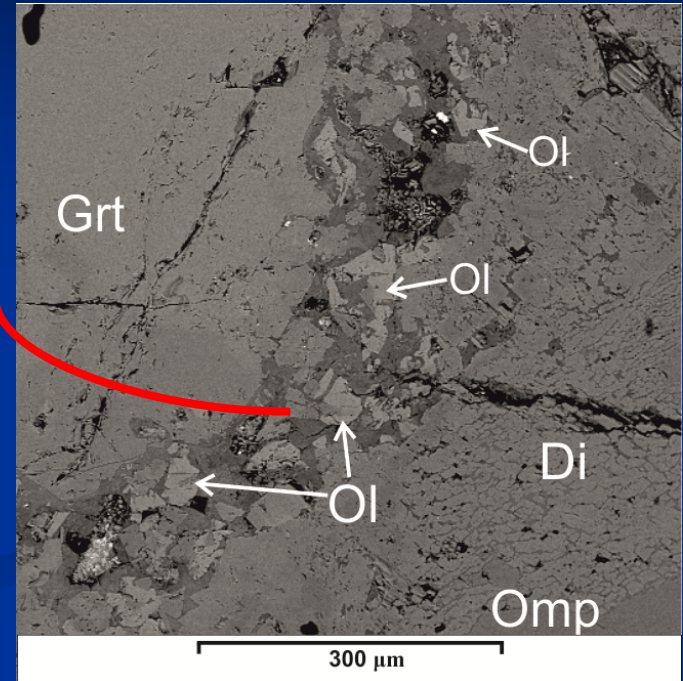
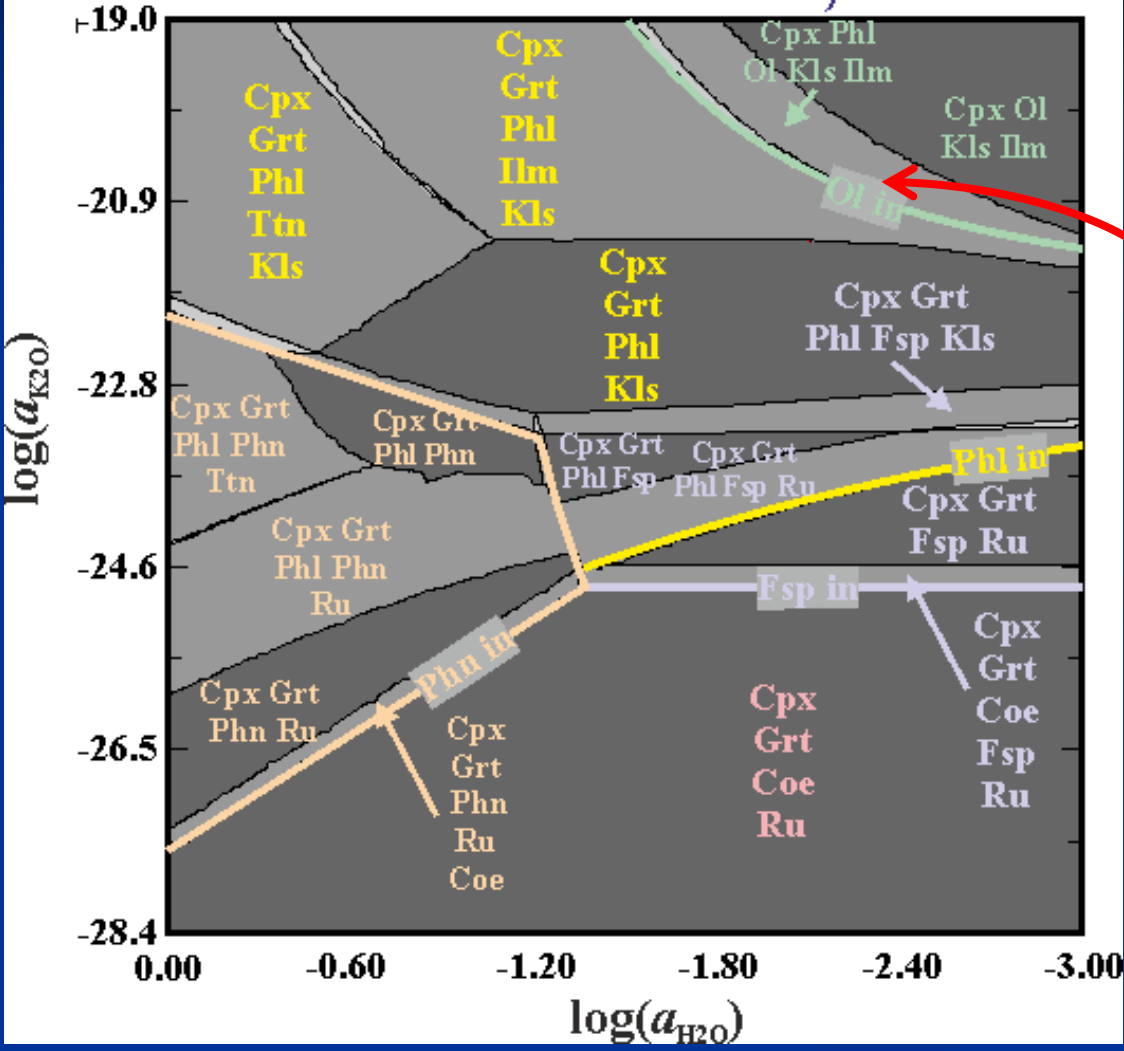


Phlogopite is a common metasomatic mineral in peridotite mantle xenoliths (Erlank et al. 1987) but not so in eclogite xenoliths, as it does not form readily by reaction of metasomatizing fluids with clinopyroxene or garnet (Kushiro and Aoki 1968)

6 GPa, 1000°C

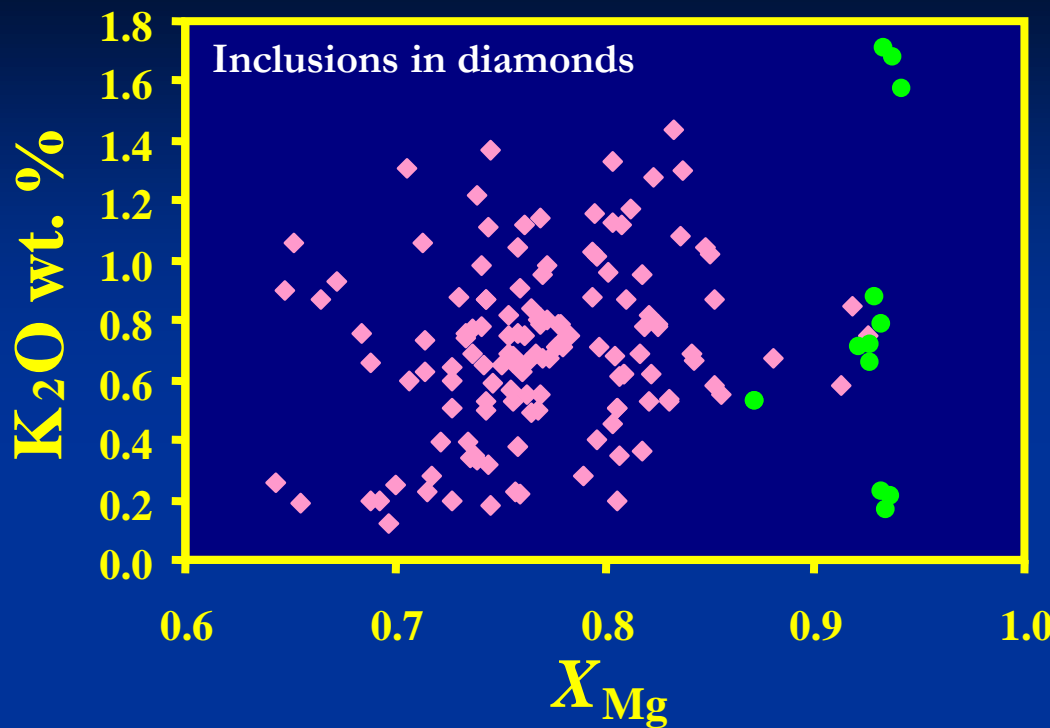


# 6 GPa, 1000°C



Eclogite xenolith, Udachnaya pipe (photo by D. Mikhailenko and A. Korsakov)

# K-bearing clinopyroxene as indicator of potassium activity



◆ Eclogite assemblages

● Peridotite assemblages

- common and stable both in peridotite and eclogite assemblages

- within a wide P-T range

- compatible with silicate and carbonate-silicate melts.

- with a potassium content regularly changing with P, T, potassium activity in a mineral forming media.

Atomistic Model of Diopside–K-jadeite ( $\text{CaMgSi}_2\text{O}_6$ – $\text{KAlSi}_2\text{O}_6$ )  
Solid Solution

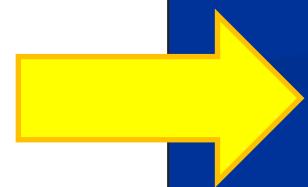
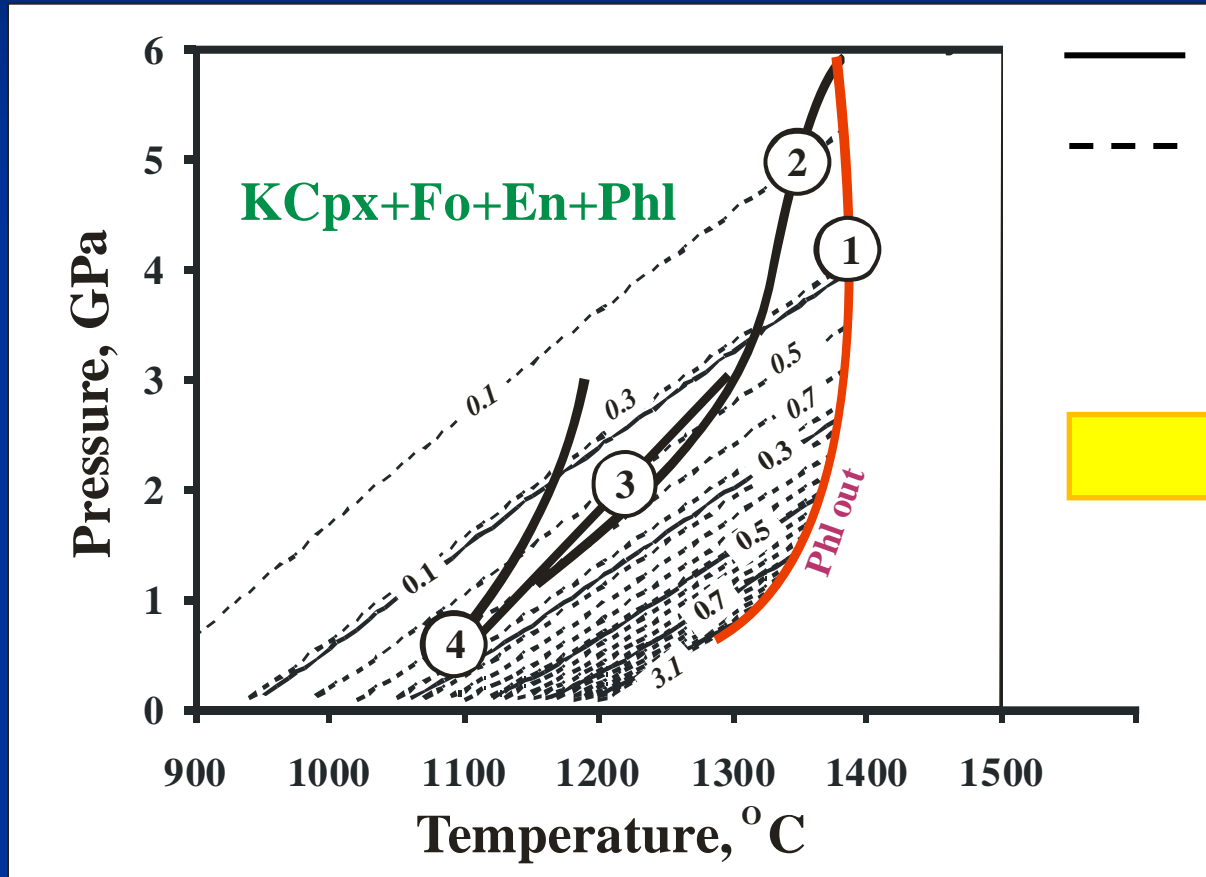
*Petrology, 2010, Vol. 18, No. 4, pp. 447–459.*

V. L. Vinograd<sup>a</sup>, O. G. Safonov<sup>b</sup>, D. J. Wilson<sup>a</sup>, L. L. Perchuk<sup>b†</sup>, L. Bindi<sup>c</sup>, J. D. Gale<sup>d</sup>, and B. Winkler<sup>a</sup>

	$\Delta_f H$ kJ/mol	$S$ J/K/mol	$V$ J/bar/mol	$K$ (GPa <sup>-1</sup> )	$\alpha$ $10^{-5}$ (K <sup>-1</sup> )
$\text{KAlSi}_2\text{O}_6$ (KJd)	-2932.7	141.24	6.479	145	3.3 (298–1298 K)

# The isopleths of KJd in the Di-KJd s.s. in the assemblage Phl+Fo+En

Reaction	$\Delta H_{298,1}$ (kJ/mol)	$\Delta S_{298,1}$ (J/mol/K)	$\Delta V_{298,1}$ (J/mol/bar)
$\text{KJd} + 2\text{Fo} + \text{H}_2\text{O} = \text{En} + \text{Phl}$	-246.36	-125.99	2.884



$a_{\text{K}20}$

- 1 – stability curve of pure Phl (Trónnes, 2002)
- 2, 3 – Di+Phl solidi (Luth, 1997; Modreski & Boettcher, 1973)
- 4 – Phl+Di+H<sub>2</sub>O solidus (Modreski & Boettcher, 1973)



*Thank you for attention*